IMPACT OF RISING SEA SURFACE TEMPERATURES ON FREQUENCY OF TROPICAL STORMS AND THEIR RELATIONSHIPS OVER NORTH INDIAN OCEAN

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ABSTRACT

Analysis of 118 years of data relating to tropical storms highlights that there is a significant change in frequency of tropical storms with rising Sea Surface Temperatures (SSTs) environment over the North Indian Ocean (NIO) on annual, seasonal and sub-seasonal scales. Decadal variability study pointed out that time-series of annual frequency of tropical storms and SST datasets reveal conspicuous turning points in the years, 1970 and 1966 respectively, which coincide with the starting of global warming period. Further there is a significant decreasing trend in the former, while a clear cut increasing tendency is observed in the SSTs field and both trends are statistically significant at 1% level over the NIO. The relationship between them is inverse and the coefficient of correlation between them is -0.52. Next, tropical severe cyclones, which appear maximum in the month of November over NIO have a direct relationship with rising SSTs and a coefficient of correlation between them is to a tune of 0.89 for the above study period. Secondly in the context of recent global warming scenario, datasets (1981-2008) relating to frequency of the tropical systems over the BOB are analyzed in Pre and Post-monsoon seasons; it is observed that there is a clear cut decreasing trend. To find out the possible reason for the decreasing trend, the authors also examined the possible relationships among frequency of the tropical systems, SSTs over the BOB, Madden & Julian Oscillation (MJO) and Southern Oscillation (SO) Indices separately. Total number of systems over the BOB is highly influenced by above SSTs only in the Pre-monsoon season, while they are significantly correlated with MJO and SO Indices in the Post-monsoon season. Above relationships are very robust in the month of November due to high frequency of tropical cyclones in every year. Finally there is a contrasting difference in the frequency of total number of tropical cyclones in the years of El Nino and La Nina episodes.

Keywords: Tropical cyclones, SSTs, MJOI, SOI.

INTRODUCTION

The Intergovernmental Panel on Climate Change has estimated that the Earth's average global surface temperature has increased by about 0.7°C in this century due to global warming, which results in change of physical characteristics of tropical oceans and genesis of storms. In 2009 a record of highest surface-airtemperature registered over India resulted in a failure of both southwest and northeast monsoon rainfall in terms of lack of rain-bearing systems over the BOB, which cross East coast of India. In this direction several other studies addressed that the SSTs over NIO is characterized by a significant climate variability and change in frequency of tropical cyclones along with other ocean basins in the tropics. Of several systems, tropical cyclones are among the most destructive natural hazards of the world; approximately 80-90% of tropical systems occur over the tropical oceans annually with a maximum peak in July and August. But in the Indian Ocean the tropical systems occur about 7% of the global frequency with primary peak in Post-monsoon (November) and secondary peak in Pre-monsoon (May) seasons (Subbaramayya and Rao, 1981 and1984; Singh et al., 2001). It is well known that the tropical cyclones are mainly triggered by both thermodynamic and dynamic parameters. In many cases, thermodynamic parameters are closely linked with each other in the tropics. Our basic understanding of tropical cyclones suggests that there could be a relationship between tropical cyclone activity and greater than SSTs of 26.5°C (Chan 1985; Gray et al., 1991; Singh and Khan, 1999; Goldenberg et al., 2001; Sujata et al., 2005). Secondly, the dynamic parameters provide necessary strength for the strengthening of tropical cyclones (Gray, 1985, 1993). Numerous studies have addressed the issue of changes in global frequency and intensity of tropical cyclone in the climate change scenario (Rajeevan, 1989; Mark et al., 1999; Mandal et al., 2007; Jayanthi, 1997; Jayanthi and Govindachari, 1999; Jayanthi and Raghavan, 2002; Chan and Shi, 1996; Chan, 2006). Later Singh et al. (2001) has reported that, 25% increase to severe cyclone stage over the NIO during November, which accounts for the highest monthly average of severe cyclone frequency. The publications about dynamical/physical relationships with global tropical cyclonic activity over Indian Ocean were carried out by Liebmann et al. (1994), Bessafi and

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Wheeler (2006) and Ho et al. (2006). Emanuel (2005) produced evidence for a substantial increase in the power of tropical cyclones during the last 50 years. This result is supported by the findings of Webster et al. (2005) who observed a substantial global increase of the most severe tropical cyclones during 1970-1995. Some more investigators have studied the changes in the tropical cyclone frequency in the Indian Ocean/BOB with a limited data (Das, 1972, 1994; IMD, 1979, 1997, 1999; Chan and Shi, 1996; Raghavendra, 1973; Ghosh, 1981; Mooley, 1980, 1981; Mooley and Mohile, 1983, 1984; Miyan, 1996; Raghavan, 1991, 1997; Oxfam India, 2000; Srivastava et al., 2000; Bhanu Kumar et al., 2004). Later there was a lot of focus on El- Nino/Southern Oscillation (ENSO) and inter-annual variability of basin wide-scale tropical cyclone activity and impact is variable by subbasin regions. For instance, ENSO is known to influence cyclone frequency in different ocean basins (Basher and Zheng, 1995; Hastings, 1990; Gupta and Muthuchami, 1991: Gupta et al., 1991). However, interest in better understanding and prediction of relatively short-term (intraseasonal and interannual) fluctuations as well as longer (decadal, multidecadal, centennial and millennial) variations has grown tremendously in the past several years in this direction. In the recent times, there has been an explosion in publications and research into how the impact of manmade global warming may be changing tropical cyclone characteristics today and decades into the future and several teleconnections are identified for possible relationships with tropical cyclones in different basins of the globe. Results of global modeling for doubled CO₂ scenarios are contradictory with simulations showing a lack of consistency in projecting an increase or decrease in the total number of tropical cyclones, although most simulations project an increase in tropical cyclone intensity (Ali, 1995; Joseph, 2004; Hulme and Viner, 1998; Lander and Guard, 1998; Walsh and Pittock, 1998; Royer et al., 1998; Landsea, 2000; Lighthill et al., 1994; Nichollas et al., 1998; Henderson et al., 1998; Tsutsui, 2002; Knutson and Tuleya, 2004; Bhanu Kumar et al., 2008). So compared with the change of frequency of tropical cyclones over other global ocean basins (Zishkha and Smith, 1980; Anyamba, 1992; Gray et al., 1991, 1992; Brian et al., 1988; Webster et al., 2005) there are a limited studies relating to changes in frequency of tropical cyclones with MJO/SSTs/ ENSO over the NIO and the BOB. Hence the aim of this paper is to investigate changing trends of tropical cyclone frequencies in both NIO and BOB based on increasing SSTs and to identify the possible relationships with global atmospheric/oceanic indices under the influence of present global warming scenario in the above study regions.

DATA AND METHODOLOGY

The main datasets used in this study are monthly frequency of tropical storms over NIO from the India

Meteorological Department (IMD) and mean monthly anomaly SSTs from NCEP/NCAR reanalysis (Extended Kaplan SSTs) over NIO for the period, 1891-2008. Frequency of tropical storms data over the BOB is obtained from the IMD, while data sets relating to MJO, SOI and anomaly SSTs over the BOB are downloaded for the period, 1981-2008. The MJO index is calculated by taking the mean values at 70°E, 80°E and 100°E in this study for both Pre (April and May) and Post-monsoon (October-December) seasons. It is interesting to note that the MJO index is varied from 0.5 (1997) to -0.7 (2008) for the month of November and they are related with the frequency of tropical storms in the same month for possible relationship. Similarly, the magnitude of SOI varied from 2.8 (1988) to -0.53 (1982) in November in this study. The Joint Typhoon Warning Centre data is used for the quality check of missing data relating to frequency of tropical storms available from the IMD. The tropical systems in the BOB play a crucial role in controlling rainfall over East coast of India during Pre and Post-monsoon seasons and a special attention is made in the present study to know the effect of global warming on frequency of tropical cyclones for the satellite era period (1981-2008). The global ocean-atmospheric indices data is downloaded from Climate Prediction Center for the period, 1981-2008. The state-of-the art of methodologies used in this study are Cramer's test for decadal variability and simple correlation and regression analyses for knowing the relationships between dependent and independent parameters.

RESULTS AND DISCUSSION

Decadal variability of frequency of tropical storms and SSTs over NIO

Global tropical SSTs have a tremendous impact on the frequency of tropical storms in the present global warming era (Emanuel, 2005; Chan, 2006). In this study the authors have examined 118 years data to examine changes in frequency of tropical storms at seasonal and sub-seasonal time scales over the NIO with rising SSTs environment using Cramer's test. Data set indicates that there are 263 numbers of storms in total (37% of the total cyclones) with a mean life period of 2.5 days, which struck the East coast of India. Figure-1a shows time-series of annual frequency of tropical storms after applying 31year Cramer's t-statistics test for climatological variability of tropical storms over NIO for the period, 1891-2008 and registered a turning point in the year 1971. It also depicts a continuous decreasing trend of frequency of tropical storms. The algorithm for the trend is represented by $F_t = -$ 0.12X+8.5 where 'F_t' is frequency of annual total number of systems and 'X' is the time. The root mean square error (RMSE; R^2) is 0.71, which is statistically significant at 0.1% level. This study is very similar to that of Webster et al. (2005) with different datasets except for different Ocean basins. Similar study is also extended using SST

El-Nino	El-Nino Pre		e May		Post		Nov	
Years = 16 La Nina = 8	El-Nino	La-Nina	El-Nino	La-Nina	El-Nino	La-Nina	El-Nino	La-Nina
Cyclones	12	4	10	2	34	18	16	8
10							F _(t) = -0.12x + 8 R ² = 0.71	.5
total number of systems								
Jo -> - 10 - -15						. III		-
1891 18	96 1901 1906 191	1 1916 1921 1926	1931 1936 1941	1946 1951 1956 :	1961 1966 1971	1976 1981 1986 1	991 1996 2001 2	2006

Table 1. Number of tropical storms/cyclones in Bay of Bengal (BOB) during El Nino and La Nina years for the period 1981-2008.

Fig. 1a. Values of Cramer's t-statistics for the 31-year running mean depicting climatological variability of annual frequency of tropical cyclones over NIO for the period, 1891-2008.

Year



Fig. 1b. Same as above except for SSTs.

field datasets over NIO and is represented by a turning point of year 1966; it also represents continuous increasing trend which is reversal to the frequency of storms over the study region (Fig. 1b). The algorithm is $I_{SSTs} = 0.12X-7.6$ with RMS error of 0.83 (0.1% level significance). This study clearly demarcates the turning point years 1971 and 1966 in the time series of frequency of tropical storms and SSTs respectively, which coincide with the recent global warming trend (1970). Increase of SSTs generally alters frequency of storms through latent and sensible heat transfer fluxes. If SSTs is greater (less) than air temperature over the NIO, it transmits (extracts) large amounts of heat to the upper atmosphere. In view of higher SSTs due to global warming in the recent four decades, more latent heat transfer takes place through evaporation. Hence higher SSTs contribute higher rate of evaporation due to higher energies of water molecules that are necessary for increase/decrease of formation of tropical storms. Of course, other important factors also equally play an important role in the genesis and modulation of tropical systems over NIO. Hence this study confirms that increase of SST field alone cannot



Fig. 2. Relationship between annual number of systems and SSTs over NIO for the period, 1891-2008.



Fig. 3. Same as above except for frequency of severe cyclonic systems and SSTs in November.

enhance the frequency of tropical cyclones. The formation of tropical storms over the NIO is due to a series of feedback mechanisms, wherein SST field is one among them.

Relationship between frequency of severe tropical cyclones and SSTs over NIO

To quantify the relationship between the above two parameters, the coefficient of correlation between them is evaluated and it is figured to a tune of -0.52, which is significant at 1% level (Fig. 2). This relationship exists in 76 years out of 118 years. It further states that the increasing SSTs result in decreasing of number of storms and vice-versa. For example in the strong El Nino year, 1998 the anomaly mean SSTs was 0.68 when the number of storms are less than two systems. This direct relationship is not yet detected in any other previous

studies so far. Thus global warming period (1970 onwards) has a tremendous impact on decreasing number of tropical storms over the NIO. Though latent and sensible heat fluxes are favorable to some extent for the formation of tropical storms by rising SSTs, the situation is reverse due to failure of other important conditions for the formation of storms like weak vertical wind shear and ocean driving wind force etc. Now this study is extended specially to severe tropical cyclones of November, which are very devastating in nature (Fig. 3). Severe cyclonic storms (>48 knots) are very destructive and dangerous coastal hazards of the NIO. The coefficient of correlation between them is positive, which amounts to 0.89 (0.1%) level significance) and it shows that the increase of SSTs enhance the frequency of severe cyclonic systems only. A diagnostic study is needed separately in this direction.



Fig.4. Scatter diagram between frequency of tropical systems over NIO in November and SSTs in October during 1891-2008.



Fig. 5. Relationship between tropical storms over BOB and MJO index in post monsoon season during, 1981-2008.

Further to predict number of severe cyclonic storms in November using October mean SSTs over NIO the relationship is examined for the study period. The relationship between them is direct (r = 0.9) and is shown in figure 4. This direct relationship indicates that increasing of SSTs infer decreasing of frequency of tropical storms over the NIO. Significant coefficient of correlation suggests an algorithm, $T_c=0.79X+0.04$ to estimate cyclones in November. The algorithm is tested for the last five years (2004-2008). Though the number of tropical cyclones varied from 0 to 4, the frequency of tropical cyclones has registered increasing trend with rising SSTs over NIO in the present study.

Frequency of tropical storms in the BOB in Pre and Post-monsoon seasons

The annual frequency of tropical systems is five times more over the BOB than over the Arabian Sea in NIO and special attention is made in this study. Authors made use of available data (1981-2008) to examine the change in frequency and trends of tropical systems separately over the BOB in the context of global warming. Analysis of annual frequency of tropical systems (5 year running average) indicates decreasing trend of total number of tropical systems over BOB, which is not significant. This is very similar to that of significant decreasing trend of frequency of tropical systems over NIO. Coming to the seasonal variability of tropical systems, they are in increasing trend during Pre-monsoon, while they are in decreasing trend in Post-monsoon seasons in the present study.

From the Indian perspective, cyclones over the BOB cause lot of damage on East Coast of India in a greater way on annual, seasonal and sub-seasonal time scales. The high mountain ranges and low-lying coastal plains and river deltas of the BOB combine to make this region extremely vulnerable to tropical systems. Authors have separately examined the impact of increasing SSTs on frequency of tropical systems over BOB in Pre-monsoon season and the relationship between them is direct, which is not statistically significant. The anomaly SSTs varied from 0.7 (1998) to -0.5 (1999) in this study. However above relationship (r = 0.42) between them in May, which

is typical month of Pre-monsoon is statistically significant at 10% level (Fig. 4). Though the SSTs over the BOB show different relationships with total number of storms and severe cyclonic storms in Pre and Post-monsoon seasons, there is a pressing need to examine how the storms are influenced by the other important factors like neighbouring coupled-ocean-atmospheric phenomina (MJOI and SOI).

Impact of MJOI and SOI on frequency of tropical systems in Post-monsoon season

Several studies reveal that tropical systems are by and large influenced by the active phases of MJO and SOI (Emanuel, 2005; Chan and Shi, 1996). In the present statistical study tropical storms were not influenced by above ocean-atmospheric predictors in the Pre-monsoon season. On the other hand, these tropical systems are partially influenced to some extent by the above predictors in the Post-monsoon season. The MJO is one of many factors that contribute to the development of tropical cyclones and it is very strong in boreal winter. Analysis of 28 year (1981-2008) data reveals that there is significant impact of MJO index on the frequency of tropical systems in BOB during Post-monsoon season. Year-to-year variations of MJO index (October-December) and frequency of tropical systems are related and figure 5 shows an inverse relationship between them (r = -0.42) and it is significant at 5% level. The above relationship holds good in 18 years out of 28 years. Further it states that the MJO index alone could not activate tropical cyclones in the BOB during Postmonsoon season.

There is no consensus among current climate models regarding how ENSO variability may change the frequency of storms (October-December) in the BOB in future. In the present study authors made an attempt to examine possible relationship between SOI and frequency of tropical storms. The maxima and minima variations in frequency of occurrence of tropical cyclones are examined in the light of El Nino and La Nina episodes. The cause of these variations appears to be non-seasonal variations in the ocean-atmosphere system. The relationship between the frequency of tropical cyclones in the Post-monsoon and SOI (October-December) is direct and it amounts to 0.40 (10% level). The above relationship holds good in 12 years out of 28 years in the Post-monsoon season. Later the impact of El Nino and its counter part La Nina on frequency of tropical storms/cyclones is also examined in detail. There are 16 El Nino and 8 La Nina episodes with 4 normal episodes during the study period. This study reveals that in the El Nino years, frequency of tropical cyclones that crossed the coast are relatively more, while in the years of La Nina the number of tropical cyclones are less (Table 1). This study supports that greater than normal rainfall in northeast monsoon season due to more number of tropical cyclones (Jayanthi and Govindachari, 2001).

CONCLUSION

Cramer's test highlights important turning points in the frequency of tropical storms and SSTs over the NIO, which coincide with the beginning of global warming era, 1970. Rising of SSTs over NIO in general and BOB in particular lead to annual decreasing of tropical storms as well as severe cyclonic storms. There is a reversal trend in the frequency of tropical storms in Pre-monsoon season and the frequencies of tropical cyclones are positively related with the SSTs over BOB.

Global atmospheric-ocean parameters close to study region show a very interesting relationship with the frequency of tropical storms over BOB. In Post-monsoon season the frequency of tropical systems are positively related with SOI, while MJO index is inversely related with the same. Above relationships are further strengthened in November which is a typical month of Post-monsoon season to cross East Coast of India.

The development of tropical storms over NIO/BOB is due to a series of feed-back mechanisms relating to ocean and atmosphere phenomena where-in MJO is one such parameter.

Thus this study also reveals the impact of El Nino and La Nina on the frequency of tropical storms/cyclones over BOB is detected.

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REFERENCES

Ali, A. 1995. Has the frequency of intense tropical cyclones increased in the NIO?.Current Science. 80(4) :575-580

Anyamba, EK. 1992. Some properties of a 20–30 day oscillation in tropical convection. J. Afr. Meteor. Soc. 1:1-19.

Basher, RE. and Zheng, X. 1995. Tropical cyclones in the Southwest Pacific: spatial patterns and relationships to the Southern Oscillation and Sea Surface Temperature. J. Climate 8. 1249-1260. Bessafi, M. and Wheeler, MC. 2006. Modulation of South Indian Ocean Tropical Cyclones by the Madden Julian Oscillation and Convectively-Coupled Equatorial Waves. Mon. Wea. Rev. 134:638-656.

Bhanu Kumar, OSRU., Krishna, KM. and SRamalingeswara, R. 2008. Simulation of environmental heavy rainfall episodes during June and July 2006 – A case study. Canadian Journal of Pure and Applied Sciences. 2(1):211-220.

Bhanu Kumar, OSRU., Naidu, CV. and S Ramalingeswara, R. 2004. Influence of Southern Oscillation and SSTs over Nino-3.4 region on the winter monsoon rainfall over coastal Andhra Pradesh. Proc. Ind. Acad. Sci. (Earth Planet. Sci.). 113(3):313-319.

Brian, RJ., Charles, J., Neumann. and Mary, ASD. 1988. A tropical cyclone data tape for the north Atlantic basin, 1886-1983: contents, limitations and uses. National Hurricane Centre, Florida (Second Printing).

Chan, JCL. 1985. Tropical cyclone activity in the northwest Pacific in relation to the El Niño/Southern Oscillation phenomenon. Mon. Wea. Rev. 113:599-606.

Chan, JCL. and Shi, J. 1996. Long-term trends and interannual variability in tropical cyclone activity over the western North Pacific. Geophys. Res. Lett. 23:2765-2767.

Chan, JCL. 2006. Comments on Changes in Tropical Cyclone Number, Duration, and Intensity in a Warming Environment. Science. 311, 1713b.

Das, PK. 1972. A prediction model for storm surges in the Bay of Bengal Nituic v 2V. 211-213.

Das, PK. 1994. Prediction of storm surges n the Bay of Bengal Proc Indian Nat Sci Ac. id v 60:513-533.

Dube, SK., Rao, AD., Sinha, PC. Murty, TS. And Bahulayan, N.1977. Mausam. 48:283-304.

Emanuel, K. 2005. Increasing destructiveness of tropical cyclones over the past 30 years. Nature. 436:686-688.

Ghosh, SK. 1981. The intensity of the Andhra cyclone of 1977. Mausam. 32:321-322.

Goldenberg, SB., Landsea, CW., Mestas-Nuñez, AM. and Gray, WM. 2001. The recent increase in Atlantic hurricane activity: Causes and implications. Science, 293:474-479.

Gray, WM. 1985. Technical Document WMO/TD No. 72, WMO, Geneva, Switzerland. (1):3-19.

Gray, WM. and Sheafer, JD. 1991. Teleconnections Linking Worldwide Climate Anomalies. Eds. Glantz, MH., Katz, MW. and Nichols, N. Cambridge University Press. 257-284.

Gray, WM., Sheaffer, JD. and Knaff, JA. 1991. Hypothesized mechanism for stratospheric QBO influence on ENSO variability. Preprints, Fifth Conf. on Climate Variations, Denver, CO, Amer. Meteor. Soc. 101-104.

Gray, WM., Landsea, CW., Mielke Jr., P. and Berry, KJ. 1992. Predicting Atlantic seasonal hurricane activity 6–11 months in advance. Wea. Forecasting. 7:440-455.

Gray, WM., Landsea, CW., Mielke Jr., P. and Berry, KJ. 1993. Predicting Atlantic basin seasonal tropical cyclone activity by 1 August. Wea. Forecasting. 8:73-86.

Gray, WM. 1993. Seasonal forecasting. Global Guide to Tropical Cyclone Forecasting, World Meteorological Organization Rep. TCP-31, WMO/TD 560. 5.1-5.21.

Gupta, A. and Muthuchami, A. 1991. El Niño and tropical storm tracks over Bay of Bengal during post monsoon season. Mausam. 42:257-260.

Gupta, GR., Desai, DS. and Biswas, NC. 1991. Cyclones and depressions over the Indian seas during 1989. Mausam. 42:1-16.

Hastings, PA. 1990. Southern Oscillations influences on tropical cyclone activity in the Australian/southwest Pacific region. Int. J. Climatol. 10:291-298.

Henderson-Sellers, A., Zhang, H., Berz, G., Emanuel, K., Gray, W., Landsea, C., Holland, G., Lighthill, J., Shieh, SL., Webster, P. and McGuffie, K. 1998. Tropical Cyclones and Global Climate Change: A Post-IPCC Assessment. Bull. Amer. Meteor. Soc. 79:19-38.

Ho, CH., Kim, JH., Jeong, H., Kim, S. and Chen, D. 2006. Variation of tropical cyclone activity in the south Indian Ocean: ENSO and MJO effects. J. Geophys. Res. 111 (22):191-210.

Hulme, M. and Viner, D. 1998. A Climate change scenario for the tropics. Climatic Change. 39:145-176.

India Meteorological Department. 1979. Tracks of storms and depressions over the Bay of Bengal and the Arabian Sea, 1877-1970. India Meteorological Department, New Delhi. pp186.

India Meteorological Department. 1997. Addendum to Tracks of Storms and Depressions in the Bay of Bengal and Arabian Sea. pp99. [Available from the Deputy Director General of Meteorology (Forecasting), India Meteorological Department, Pune-411005, India].

India Meteorological Department. 1999. Report on Cyclonic Disturbances over North Indian Ocean during 1998. Regional Specialized Meteorological Centre for Tropical Cyclones (RSMC). pp71. [Available from India Meteorological Department, New Delhi 110003, India].

Joseph, PV. 2004. Paper presented in the workshop on Global Climate Change and Tropical Cyclones, Dhaka, Bangladesh. Climate Research. 27:77-83 Jayanthi, N. 1997. An objective analysis of tropical cyclones of the North Indian Ocean with special reference to track prediction capabilities over the Bay of Bengal and Arabian Sea. Ph.D thesis, University of Madras. pp202.

Jayanthi, N. and Govindachari, S. 1999. El Niño and northeast monsoon rainfall over Tamilnadu, *Mausam*. 50, 2, 217.

Jayanthi, N and Raghavan, S. 2002. Study of Bay of Bengal cyclones using buoy data. Curr. Sci. 82:382.

Knutson, TR. and Tuleya, RE. 2004. Impact of CO_2 -Induced Warming on Simulated Hurricane Intensity and Precipitation: Sensitivity to the Choice of Climate Model and Convective Parameterization. J. Climat. 17:3477-3495.

Landsea, CW. 2000. Climate variability of tropical cyclones-past present and future. Storms. (vol. 1). Eds. Pielke, R Sr. and Pielke, R Jr. Routledge. 221-241.

Lander, MA. and Guard, CP. 1998. A look at global tropical cyclone activity in 1995: contrasting high Atlantic activity with low activity in the basins. Mon. Wea. Rev. 126:1163-1173.

Liebmann, B., Hendon, HH. and Glick, JD. 1994. The relationship between tropical cyclones of the western Pacific and Indian oceans and the Madden-Julian oscillation. J. Meteor. Soc. Japan. 72:401-412.

Lighthill, J., Holland, G., Gray, W., Landsea, C., Craig, G., Evans, J., Kurihara, Y. and Guard, C. 1994. Global climate change and tropical cyclones. Bull. Amer. Meteor. Soc. 75:2147-2157.

Mark, RJ., Beenay, P. and Bhawoodien, P. 1999. Climatic determinants and statistical prediction of tropical cyclone days in the southwest Indian Ocean. Journal of Climate. 12:1738-1746.

Mandal, M., Mohanty, UC., Sinha, P. and Ali, MM. 2007. Impact of sea surface temperature in modulating movement and intensity of tropical cyclones. Natural Hazards. 41(3):413-427.

Miyan, A. 1996. Assessment of vulnerability of people, livestock, agriculture and industry to tropical cyclones, storm surges and floods. Proc. Seminar on Meteorological and Hydrological Risk Assessment, New Delhi, India, World Meteorological Organization Rep. TCP-40, WMO/TD 761:75-119.

Mooley, DA. 1980. Severe cyclonic storms in the Bay of Bengal, 1877-1977. Mon.Wea. Rev. 108:1647-1655.

Mooley, DA. 1981. Increase in annual frequency of the severe cyclonic storms of the Bay after 1964-possible causes. Mausam. 32:35-40.

Mooley, DA. and Mohile, CM. 1983. A study of cyclonic storms incident in the different sections on the coast around Bay of Bengal. Mausam. 34:139-152.

Mooley, DA. and Mohile, CM.1984. Cyclonic storms of the Arabian Sea, 1877–1980. Mausam. 35:127-134.

Nicholls, N., Landsea, CW. and Gill, J. 1998. Recent trends in Australian region tropical cyclone activity. Meteor. Atmos. Phys. 65:197-205.

Oxfam India. 2000. Cyclones in Andhra Pradesh. Oxfam and Andhra Pradesh Government Rep. [Available from Oxfam India, 3/19, Shiva Arun Colony, WestMarredpally, Secunderabad 500 026, India].

Ragavendra, VK. 1973. A statistical analysis of the number of tropical storms and depressions in the Bay of Bengal during 1890-1969., Indian J. Meteor. Geophys. 24:125-130.

Raghavan, S. 1991. The May 1990 cyclone and its predecessors. J. Appl. Hydrol., IV, 1–8. [Available from Association of Hydrologists of India, c/o Dept. of Geophysics, Andhra University, Visakhapatnam 530003, Andhra Pradesh, India].

Raghavan, S. 1997. Radar observations of tropical cyclones over the Indian seas. Mausam. 48: 169-188.

Rajeevan, M. 1989. Post monsoon tropical cyclone activity in the North Indian Ocean in relation to the El Niño/Southern Oscillation phenomenon. Mausam. 40:43-46.

Royer, JF., Chauvin, F., Timbal, B., Araspin, P. and Grimal, D. 1998. A GCM study of the impact of greenhouse gas increase on the frequency of occurrence of tropical cyclones. Climate Change. 38:307-343.

Singh, OP. and Khan, TMA. 1999. Changes in the frequencies of cyclonic storms and depressions over the Bay of Bengal and the Arabian Sea. SMRC Rep. 2. 121.

Singh, OP., Tariq Masood, Khan, A. and Sazedur, RM. 2001. Has the frequency of intense tropical cyclones increased in the north Indian Ocean? Current Science. 80(4): 25.

Srivastava, AK., Sinha Ray, KC. and De, US. 2000. Trends in the frequency of cyclonic disturbances and their intensification over Indian seas. Mausam. 51:113-118.

Subbramayya, I. and Rao, SRM. 1984. Notes and Correspondence on frequency of Bay Of Bengal Cyclones in the Post-monsoon season, Monthly Weather review. 112:1640-1642.

Subbramayya, I. and Subba, RM. 1981. Cyclone climatology of North Indian Seas. Indian J. Mar. Sci. 10:366-368.

Sujata Mandke, K., Sahai, AK. and Shinde, MA. 2005. Dynamical ensemble seasonal forecast experiments of recent Indian summer monsoons: An assessment using new approach. IITM research report. RR-107.

Tsutsui, J. 2002. Implications of anthropogenic climate change for tropical cyclone activity: A case study with the NCAR CCM2. J. Meteor. Soc. Japan. 80:45-65.

Walsh, K. and Pittock, AB. 1998. Potential changes in tropical storms, hurricanes and extreme rainfall events as a result of climate change. Climate Change. 39:199-213.

Webster, PJ., Holland, GJ., Curry, JA and Chang, HR. 2005. Changes in tropical cyclone number, duration and intensity in a warming environment. Science. 309:1844-1846.

Zishkha, KM. and Smith, PJ. 1980. The climatology of cyclones and anticyclones over North America and surrounding ocean environs for January and July, 1950-77. Mon.Wea. Rev. 108:387-401.

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